



# Recent upgrades of operational storm surge models in Japan Meteorological Agency

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**Japan Meteorological Agency**

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# Introduction

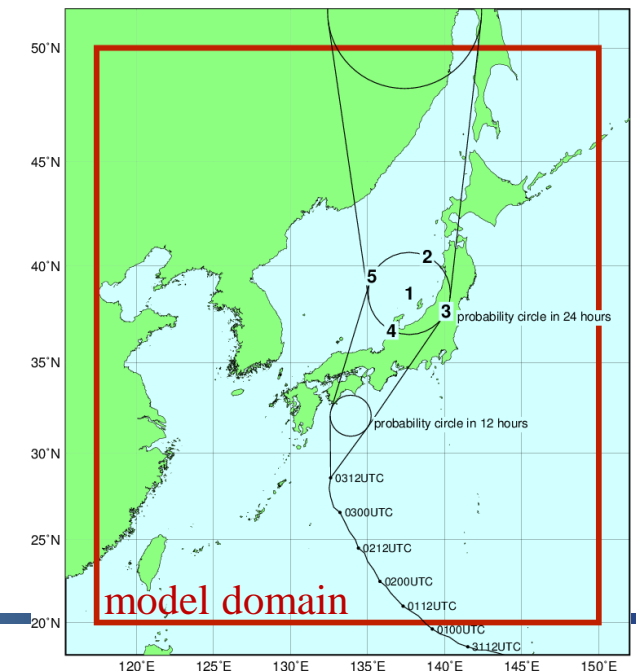
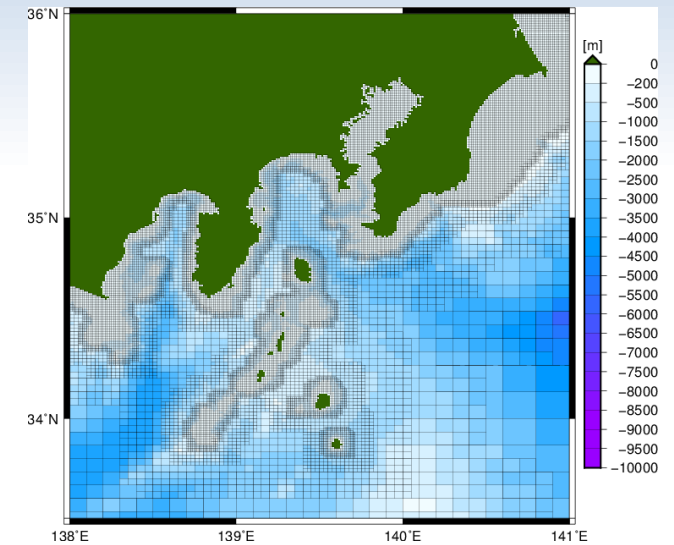
- Japan Meteorological Agency (JMA) operates two types of regional storm surge model covering the coastal area of Japan and Asia.
- Both models were based on multi-scenario predictions only for six typhoon tracks and could not provide probabilistic forecast products before 2022.  
*critical to capture the risk of storm surges within sufficient lead time.*
- JMA newly introduced storm surge ensemble prediction systems (EPSs) and started providing probabilistic forecast products from August 2022.

# Specification

Model	Two-dimensional nonlinear model
Model domain	20°N - 50°N, 117.4°E - 150°E
Grid resolution	1km ~ 16km (Adaptive Mesh Refinement)
Forecast period	78 hours (00, 12UTC) 39 hours (03, 06, 09, 15, 18, 21UTC)
Atmospheric forcing	JMA Meso-Scale Model (MSM, resolution ~ 5km) Typhoon bogusing
Member	No-typhoon case: 1 member (MSM) Typhoon case: 6 members (MSM + typhoon bogusing)
Purpose of operation	Issuing of storm surge advisories/warnings in the Japan area

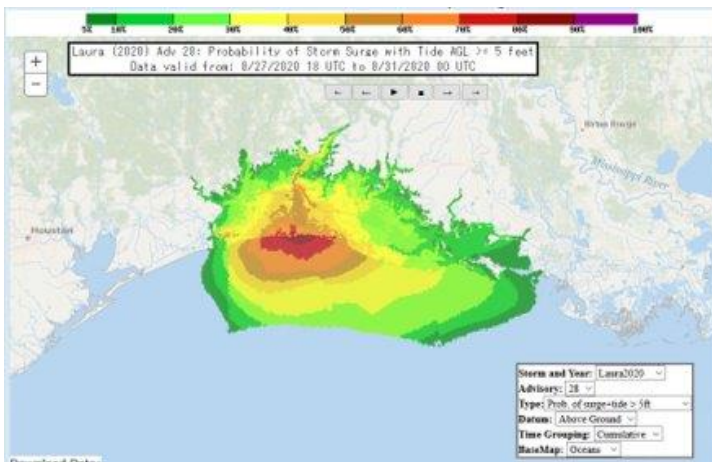
- ✓ If one or more typhoons exist near the coastal area of Japan, additional model runs are executed for five hypothetical tracks to consider typhoon forecast errors.
- Five tracks are prescribed at the center and four points on the probability circles.
  - Storm surges are calculated by implanting typhoon bogusing into each track.
- Six scenarios are considered.

Example of a grid around Japan's Kanto area



# Probabilistic Forecast System (PFS)

- Six scenarios are insufficient to consider typhoon forecast errors and capture potential storm surges.
  - A probabilistic approach is suitable for storm surge forecast with a lead time of 3-5 days.
- JMA newly introduced the Japan-area storm surge probabilistic forecast system (PFS).
  - The concept is based on the P-Surge model operated by NOAA/NHC
  - This system consists of an ensemble of the Japan-area model runs.

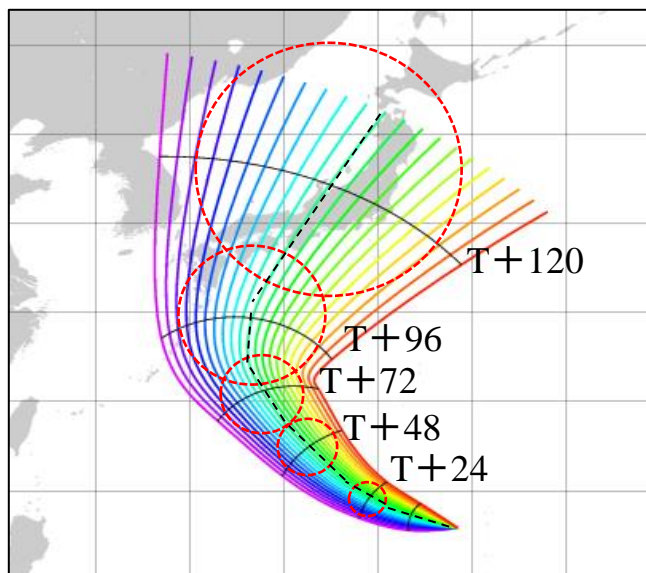


<https://slosh.nws.noaa.gov/psurge2.0/index.php>

# Method of PFS

- Create 21 cross-track members by shifting an official predicted track.
  - The whole tracks cover 70% probability circle + margin region.
  - The distance between members is the same. = inside of  $1.25 \times$  radius of 70% probability circle
  - For each track, storm surges are calculated up to 132-h by implanting typhoon bogusing.

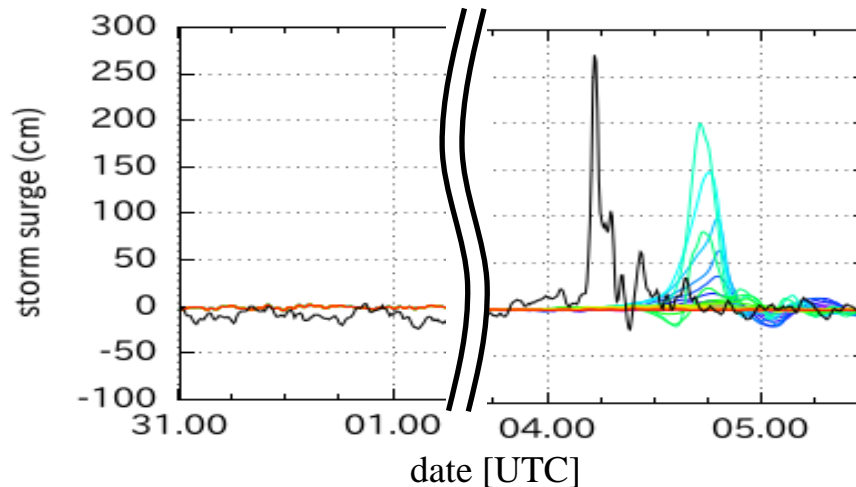
21 tracks for TY JEBI (31 Aug. 2018 00UTC Initial)



- ✓ Black dotted line: official predicted track
- ✓ Red dashed circles: 70% probability circle

Calculated 21 storm surges at Osaka (31 Aug. 2018 00UTC Initial)

- four days earlier of TY JEBI approaching



- ✓ Black solid line: observational data

- It is insufficient to consider only cross-track errors.
  - underestimate a maximum value
  - miss a timing of peak

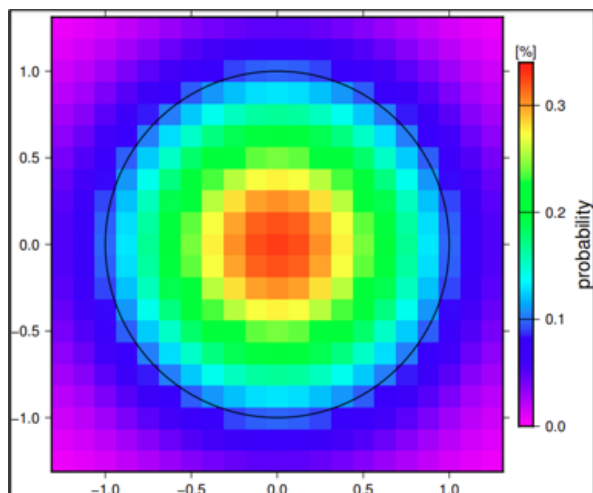


# Probabilistic forecast products

- Assign a probability to each member and calculate probabilistic quantities.
  - 2D normal distribution is assumed as the cross- and along-track error distributions related to JMA's typhoon forecast. (consistent with Best Track Data of various typhoon)

## 2D normal distribution in the PFS

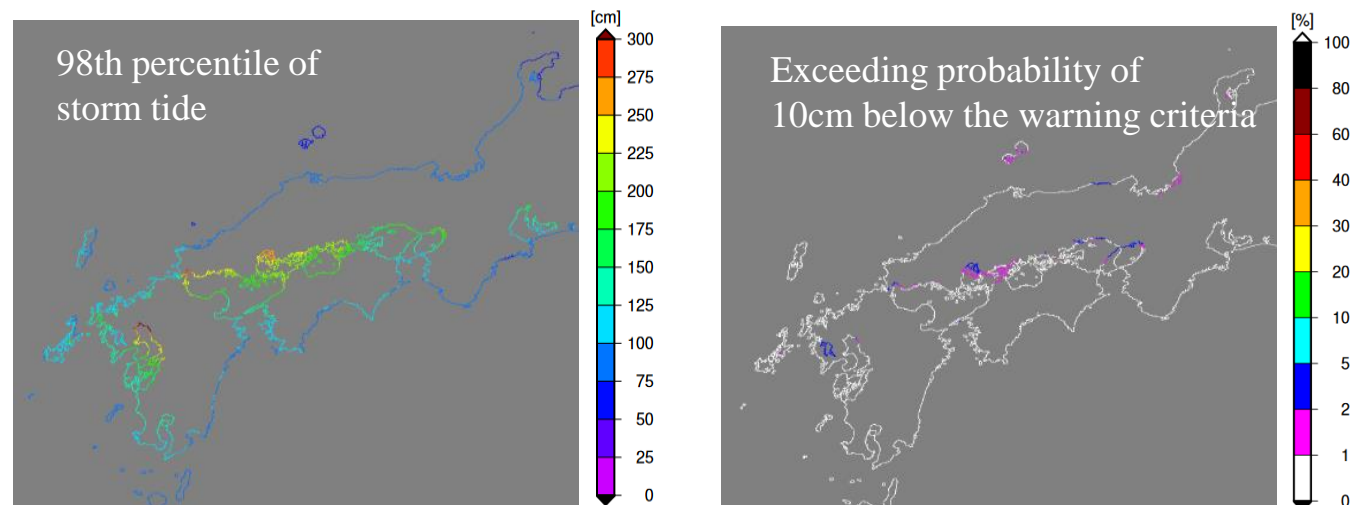
- Its origin is located at the center of typhoon.



- Black solid line: probability circle.
- The horizontal and vertical axes are normalized by the radius of a probability circle.

## Example of products in operation (31 Aug. 2018 00UTC Initial)

- four days earlier of TY JEBI approaching



- The potential storm surge can be captured in the early stage of typhoon approaching.  
→ lead to the early issue of storm surge advisories/warnings

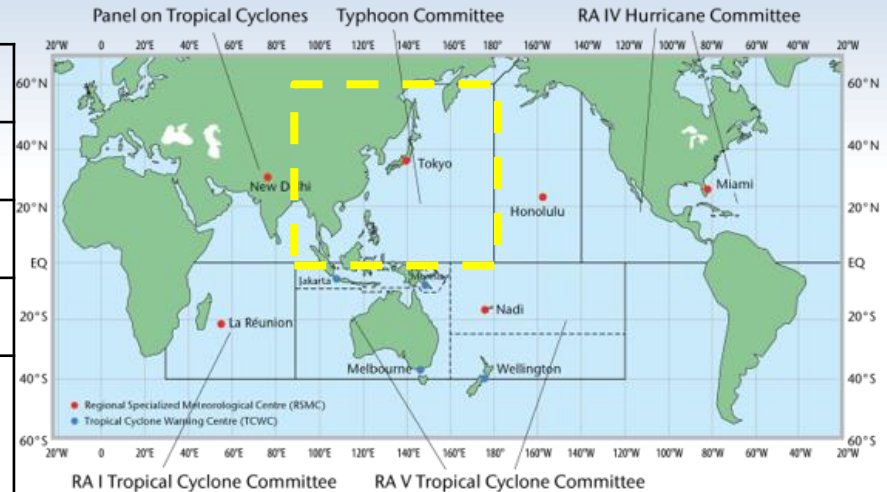
# Specification of previous model

Model	Two-dimensional linear model
Model domain	0° - 46°N, 95°E - 160°E
Grid resolution	3.7km (Lat-Lon, Arakawa C-grid)
Forecast period	72 hours (00, 06, 12 18UTC)
Atmospheric forcing	JMA Global Spectral Model (GSM, resolution ~ 13km) JMA Global Ensemble Prediction System (GEPS, resolution ~ 27km) Typhoon bogusing
Member	No-typhoon case : 1 member (GSM) <b>Typhoon case: 6 members (GSM + 5 members chosen from GEPS)</b>
Purpose of operation	Provision of storm surge prediction products to Typhoon Committee Members in the framework of WMO Storm Surge Watch Scheme

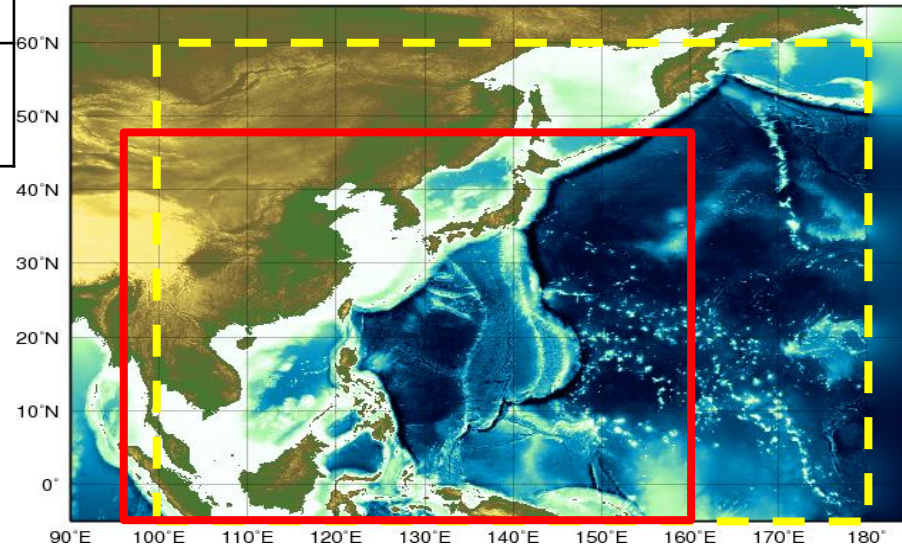
## ➤ To be improved

- Grid resolution was low in the view of storm surge predictions.
- Probabilistic forecast products could not be provided.
- Model domain did not cover Marshall Islands.

— Model domain  
- - RSMC Tokyo area



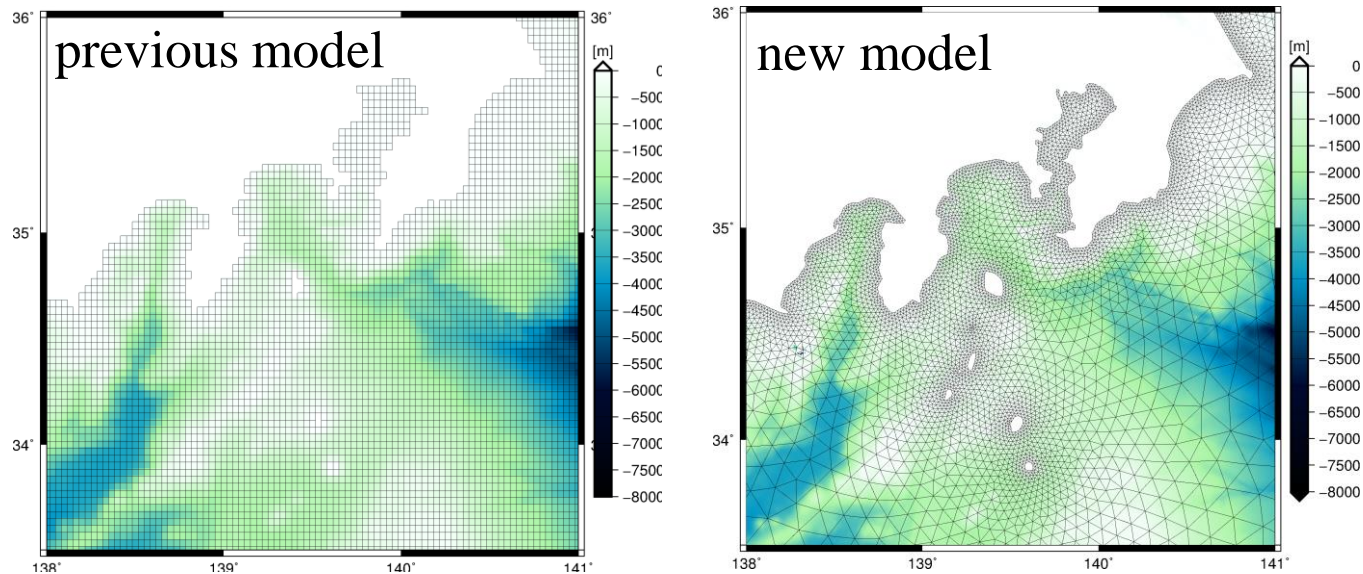
The yellow box is the RSMC Tokyo's area of responsibility  
<https://community.wmo.int/latest-advisories-rsmcs-and-tcwcs>





# Unstructured grid

- JMA has constructed its own unstructured grid finite volume model.
  - The resolution in coastal areas has been increased from 3.7km to 1.5km.
  - The grid becomes coarse with distance from coastal line (maximally around 50km).



- The expression of complex topography and small islands has been largely improved.
- The whole grid points have been drastically reduced.  
→ Computationally efficient
- The spherical non-linear 2D shallow water equation is solved by a finite volume method.

## Comparison of grid around Japan's Kanto area

- JIGSAW (Engwirda, 2017) is used as a mesh-generating tool.

cf. FVCOM (Chen et al, 2013)

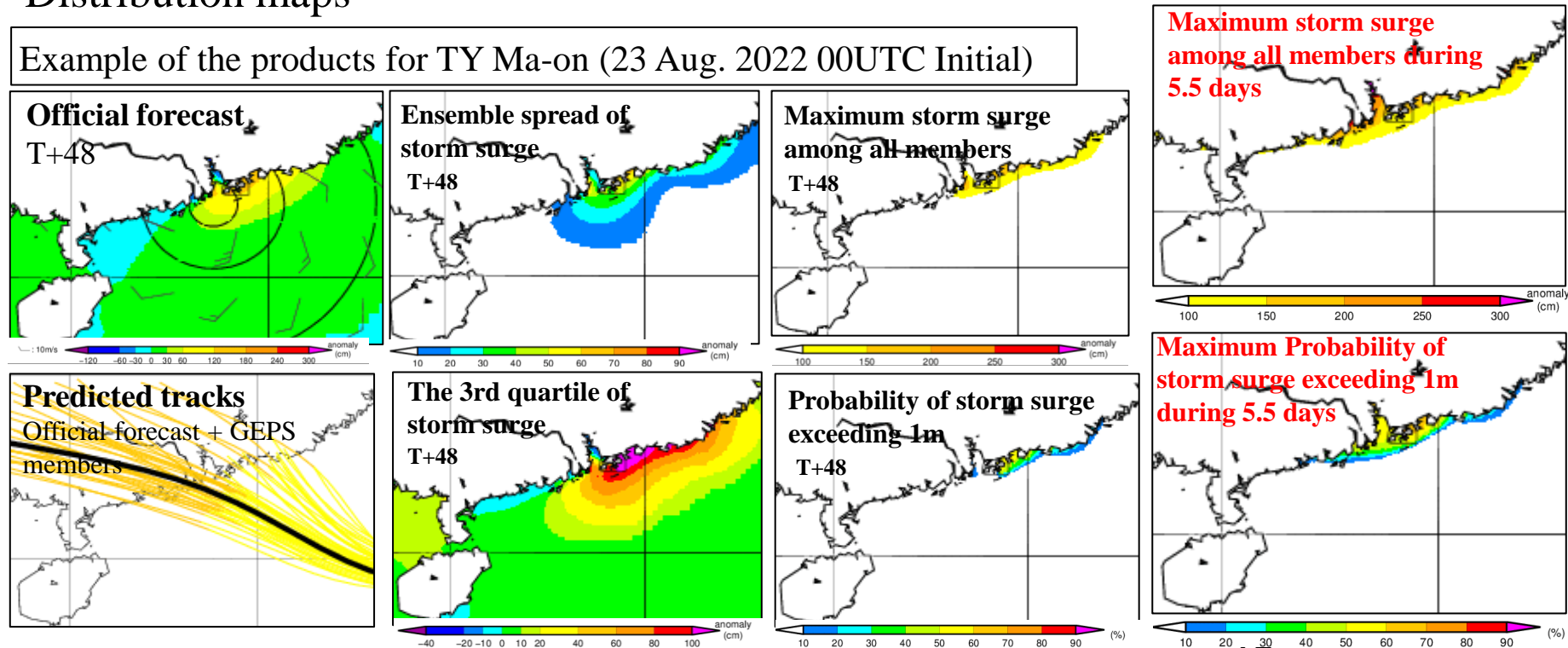
FESOM2 (Danilov et al, 2017)

# Probabilistic forecast products

- The Asia-area EPS has been built by using **GSM (1 member)** and **GEPS (51 members)**.
  - Products are provided to Typhoon Committee member countries.

## Distribution maps

Example of the products for TY Ma-on (23 Aug. 2022 00UTC Initial)

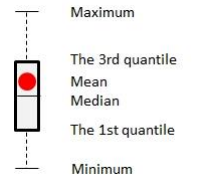


# Probabilistic forecast products

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## Time-series chart for selected locations

Example of the products for TY Ma-on (23 Aug. 2022 00UTC Initial)

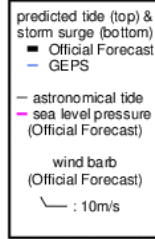
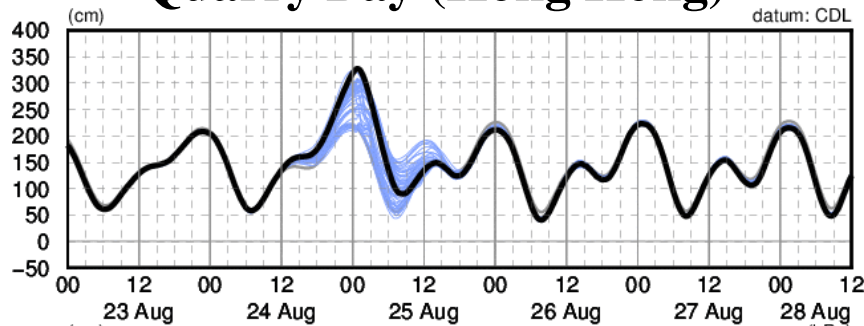


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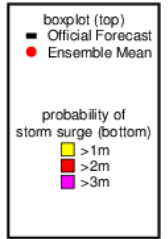
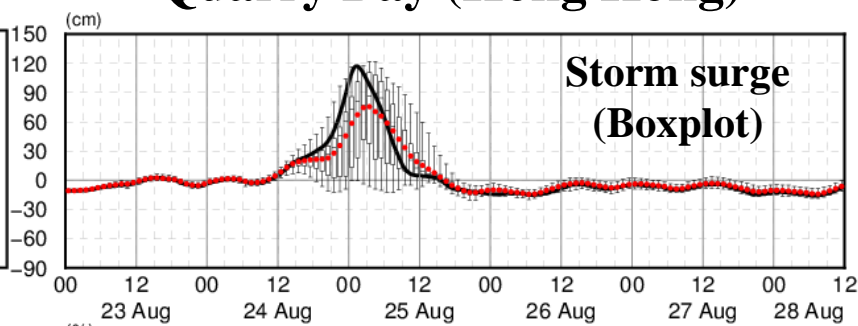
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Storm tide

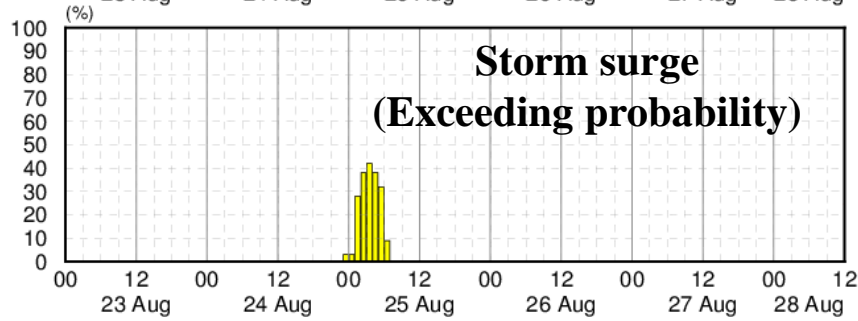
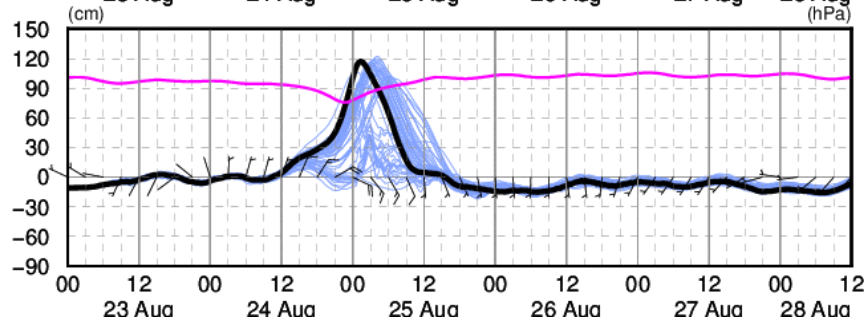
### Quarry Bay (Hong Kong)



### Quarry Bay (Hong Kong)



Storm surge & Sea level pressure



# Other upgrades

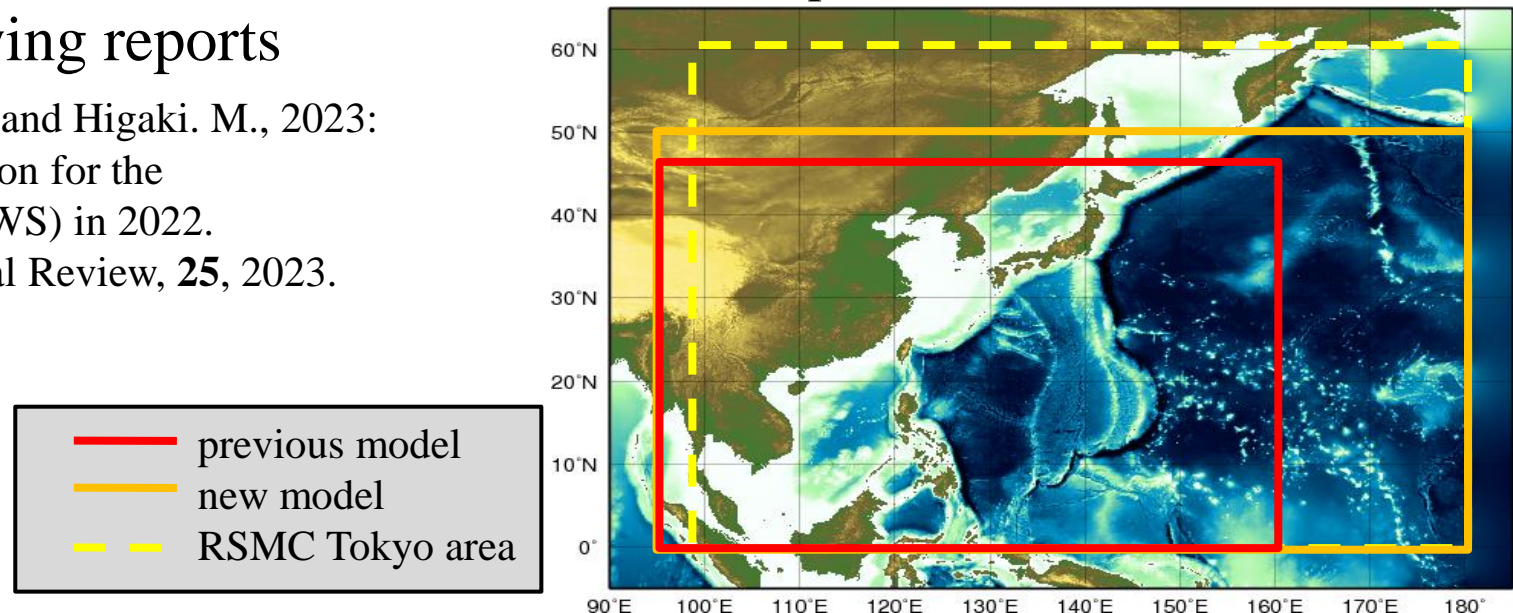
- Expansion of model domain
  - The new model can cover most of RSMC Tokyo area of responsibility (including Marshall Islands)
- Extension of forecast period from 72 to 132 hours
  - The new model yield more sufficient lead time for issuing warnings & evacuation

- For more detail, see the following reports

- ✓ Hasegawa. H., Sugano. J., Fukuura. T. and Higaki. M., 2023:  
Upgrade of JMA's Storm Surge Prediction for the  
WMO Storm Surge Watch Scheme (SSWS) in 2022.  
RSMC Tokyo-Typhoon Center Technical Review, **25**, 2023.

[Download PDF](#)

Comparison of the model domain



# Summary

- JMA extensively upgraded two types of operational storm surge models in 2022.
  
- Japan area
  - JMA started operating Japan-area storm surge probabilistic forecast system (PFS) and providing probabilistic forecast products.
  - This system is effective to capture potential storm surges within sufficient lead time when typhoons are approaching.
  
- Asia area
  - JMA introduced an unstructured grid finite volume model and built ensemble prediction system and started providing probabilistic forecast products for Typhoon Committee Members.

# Perspective

## ➤ Japan area

- introduction of an unstructured grid finite volume model
- improvement of typhoon bogusing by using a vortex relocation scheme
- incorporation of a perturbation of typhoon strength to the PFS

## ➤ Asia area

- improvement of typhoon bogusing by using a vortex relocation scheme

A vortex relocation scheme:

- Kurihara. Y., Bender. M. A. and Ross. R. J. Ross. 1993: An Initialization Scheme of Hurricane Models by Vortex Specification. *Mon. Wea. Rev.* **121** (7): 2030–45.
- Kurihara. Y. Bender. M. A., Tuleya. R. E. and Ross. J. R. 1995: Improvements in the GFDL Hurricane Prediction System. *Mon. Wea. Rev.* **123** (9): 2791–2801.

# Backup

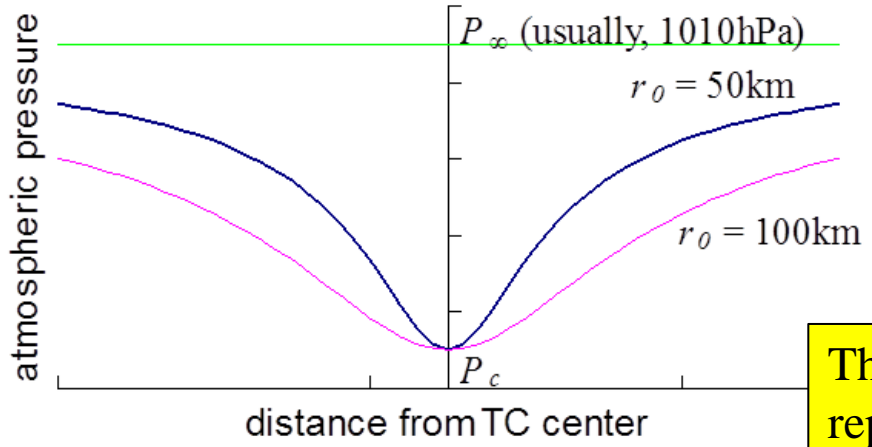
# Typhoon bogusing

Bogus is simple parametric TC model. Radial pressure distribution is represented by Fujita's formula:

$$P = P_{\infty} - \frac{P_{\infty} - P_c}{\sqrt{1 + (r/r_0)^2}}$$

$P$ : atmospheric pressure  
 $P_{\infty}$ : environmental pressure  
 $P_c$ : central pressure  
 $r$ : distance from center  
 $r_0$ : scaling factor

$r_0$  decides sharpness of pressure distribution.  $r_0$  is calculated from 30/50knot radius.



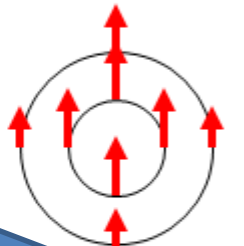
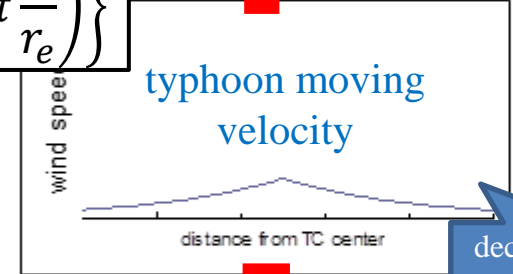
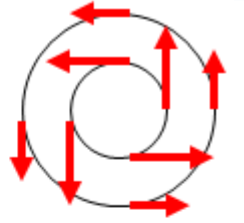
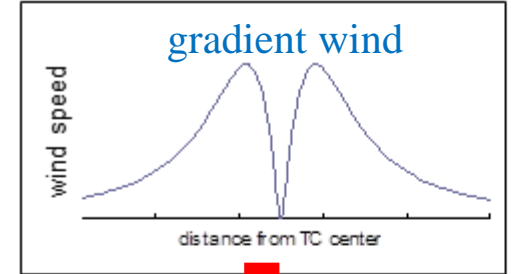
gradient wind relation:

$$-\frac{v_g^2}{r} - f v_g = -\frac{1}{\rho} \frac{\partial P}{\partial r}$$

$f$ : coriolis force  
 $\rho$ : density of air  
 $v_g$ : gradient wind

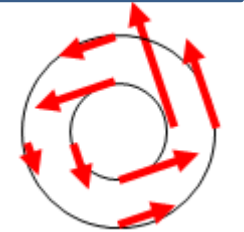
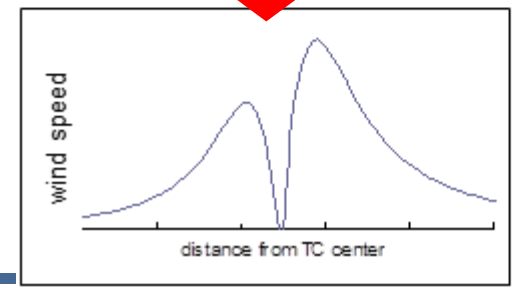
$$\mathbf{w} = C_1 \left\{ \mathbf{v}_g + \mathbf{C} \cdot \exp\left(-\pi \frac{r}{r_e}\right) \right\}$$

$\mathbf{w}$ : wind field  
 $\mathbf{C}$ : TC velocity vector  
 $r_e$ : coefficient of decay



decays exponentially with distance from the center

inflow angle →



This parametric TC model cannot represent wind speed reduction due to topography.



# Typhoon bogusing

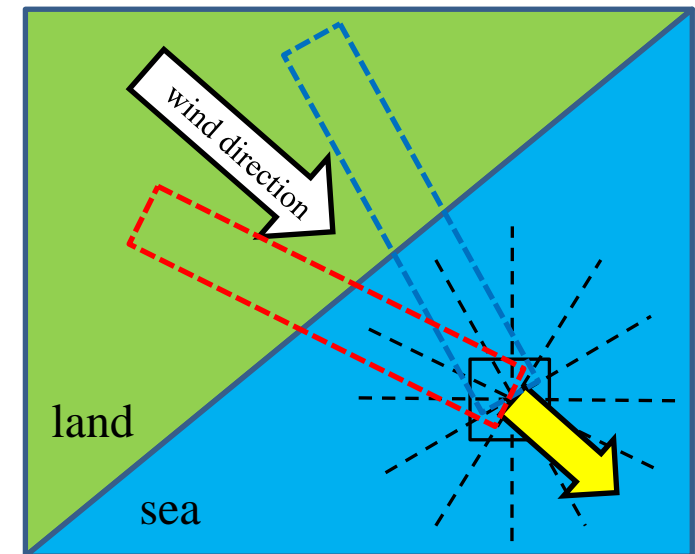
- **Upwind directional land roughness parameterization** (Westerink *et al*, 2008) is adopted to represent the wind speed reduction in coastal areas.
- Calculation of dumping coefficient

1. Calculate weighted average of roughness length of 30 km-upwind for 12-directions

$$\text{Gaussian weight : } w(i) = \frac{1}{\sqrt{2\pi\sigma}} e^{[-d(i)^2/2\sigma^2]}$$

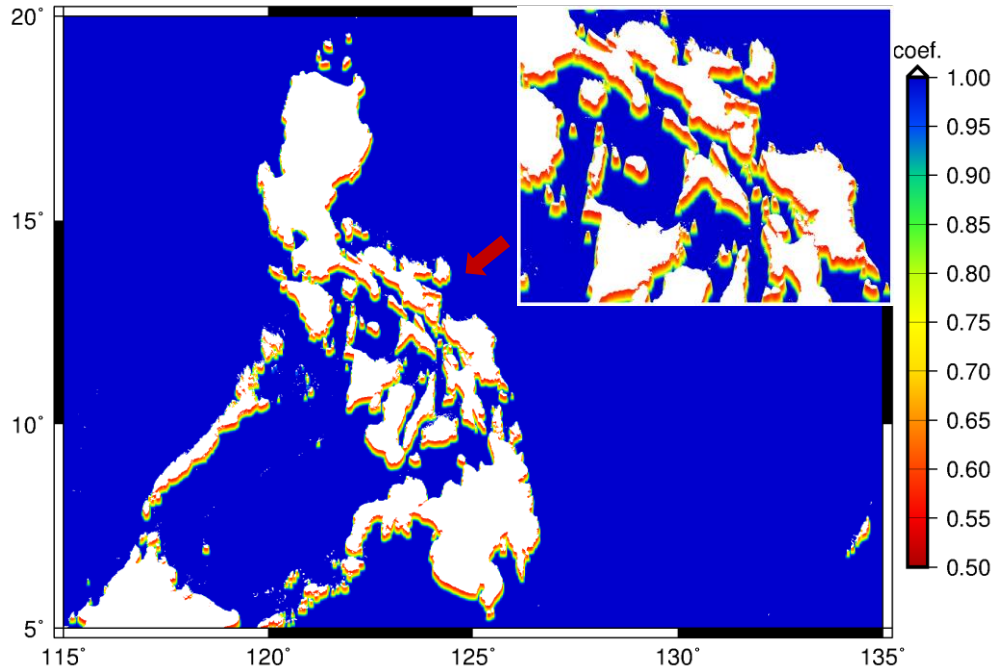
2. Interpolate directional land roughness to wind direction in a target grid
3. Calculate dumping coefficient from ratio between the interpolated land roughness and sea roughness

$$f_{\text{r-directional-k}} = \left( \frac{z_{0\text{marine}}}{z'_{0\text{land-directional-k}}} \right)^{0.0706}$$

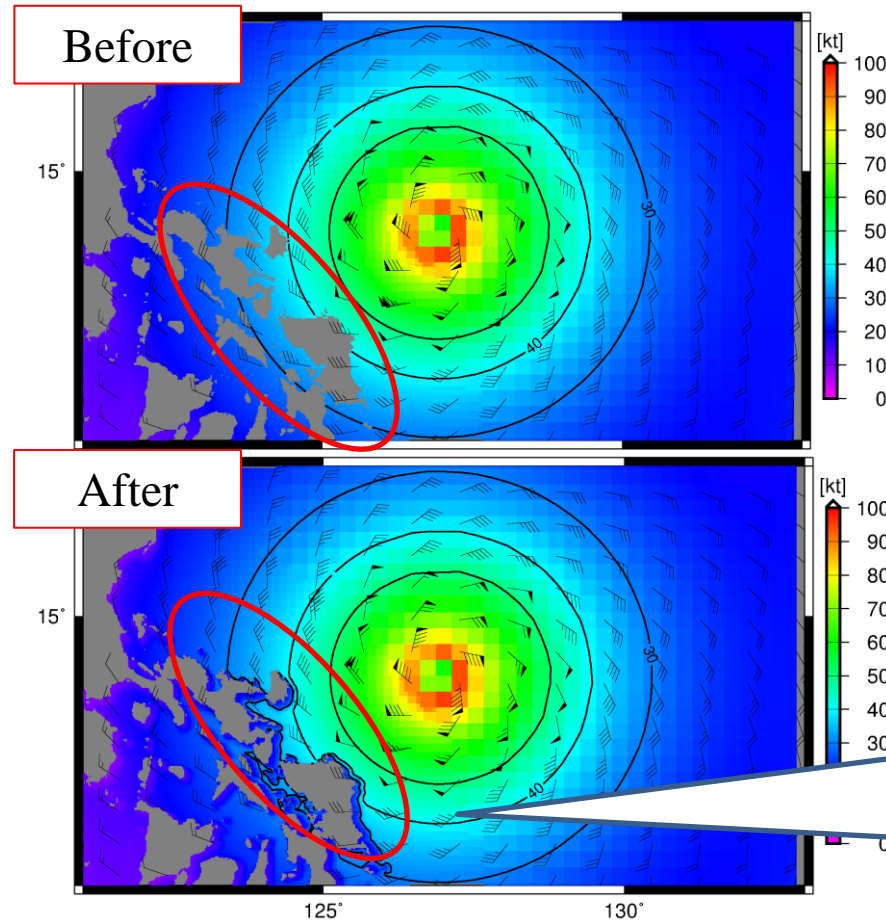


# Typhoon bogusing

➤ Effect of upwind directional land roughness parameterization



Dumping coefficient against northern wind



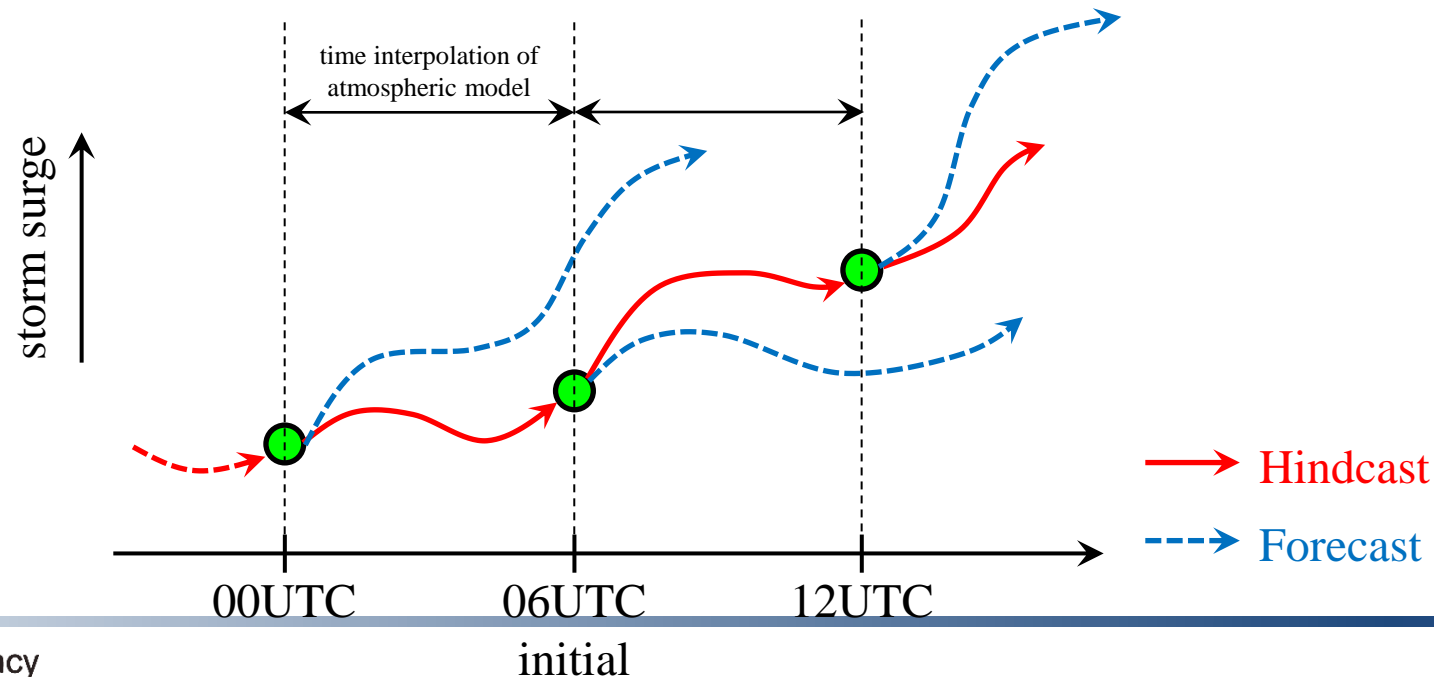
Ex.) TY Surigae  
Colors and contours show wind speed.

Offshore winds are effectively dumped.

# Initial condition

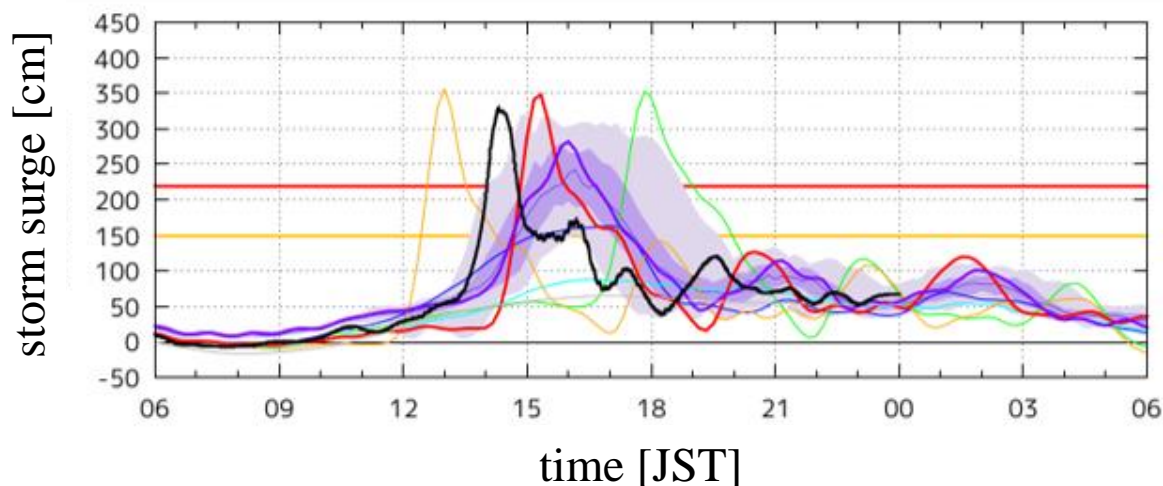
- Initial conditions are generated by “Hindcast”.
  - Initials of atmospheric model include analysis.
  - Storm surge model is driven by **previous and current initials (analysis values) of atmospheric model** (more accurate and continuous calculations).
  - No data assimilations

Accuracy of the storm surge model largely depend on the accuracies of atmospheric model.



# Storm surge EPS based on MEPS

- JMA initially considered to introduce a storm surge EPS based on JMA Meso-scale EPS (MEPS).
  - MEPS is an ensemble run of JMA Meso-scale model (resolution: 5km, the number of members: 21)
- In the performance test of storm surge EPS based on the MEPS, it was found that the worst scenario can be missed when an ensemble spread of MEPS is small.



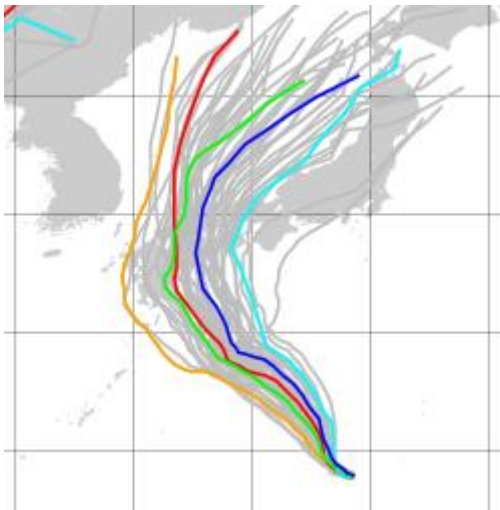
- Predicted storm surge at Osaka (03 Sep. 2018 06UTC Initial)
  - Dark purple band: 1st - 3rd quartile
  - Thin purple band: area between maximum and minimum value
  - Purple line: ensemble mean
  - Black line: observational data

# Cluster analysis

➤ In the previous model, five members was extracted from GEPS by a cluster analysis.

## ➤ K-mean method

- All members are classified into five clusters according to similarity of tracks.
- The nearest members are chosen from each center of clusters.



### K-means method

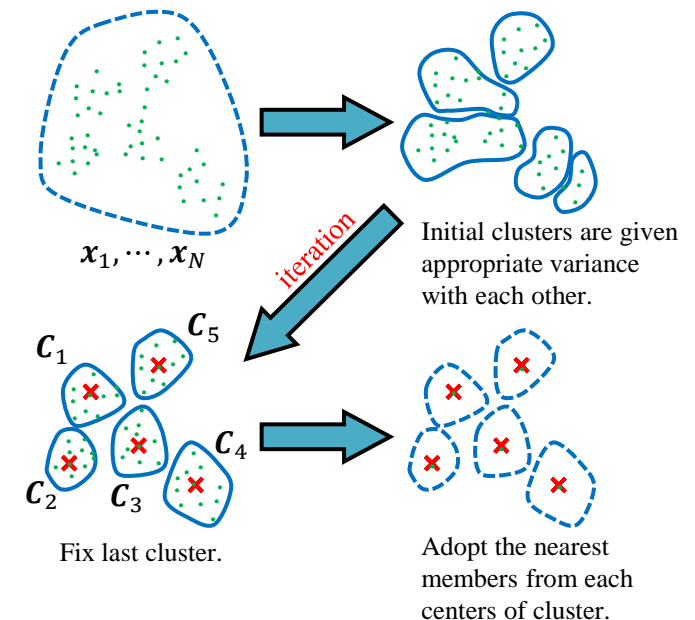
(N=51, K=5)

Center of typhoon:

$$\mathbf{x}_i = (\text{lat}_i, \text{lon}_i), (i = 1, \dots, N)$$

Center of cluster:

$$\mathbf{c}_k = \frac{1}{N_k} \sum \mathbf{x}_i, (k = 1, \dots, K)$$



➤ Example of TY Krosa

Gray line: Tracks of GEPS for 51 members.

Color line: Five members extracted by a cluster analysis

(Hasegawa, Kohno, Higaki and Itoh, 2017)

# Finite volume model

cf. FVCOM (Chen et al, 2013), FESOM2 (Danilov et al, 2017)

## ➤ Solve 2d shallow water equation in spherical coordinate system

momentum equation

$$\frac{\partial U}{\partial t} + \frac{1}{r \cos \varphi} \left( \frac{\partial u^2 H}{\partial \lambda} + \frac{\partial uvH \cos \varphi}{\partial \varphi} \right) - fvH - \frac{uvH}{r} \tan \varphi = -\frac{gH}{r \cos \varphi} \frac{\partial(\zeta - \zeta_0)}{\partial \lambda} + \frac{\tau_{s\lambda}}{\rho_w} - \frac{\tau_{b\lambda}}{\rho_w}$$

$$\frac{\partial V}{\partial t} + \frac{1}{r \cos \varphi} \left( \frac{\partial uvH}{\partial \lambda} + \frac{\partial v^2 H \cos \varphi}{\partial \varphi} \right) + fuH + \frac{u^2 H}{r} \tan \varphi = -\frac{gH}{r} \frac{\partial(\zeta - \zeta_0)}{\partial \varphi} + \frac{\tau_{s\varphi}}{\rho_w} - \frac{\tau_{b\varphi}}{\rho_w}$$

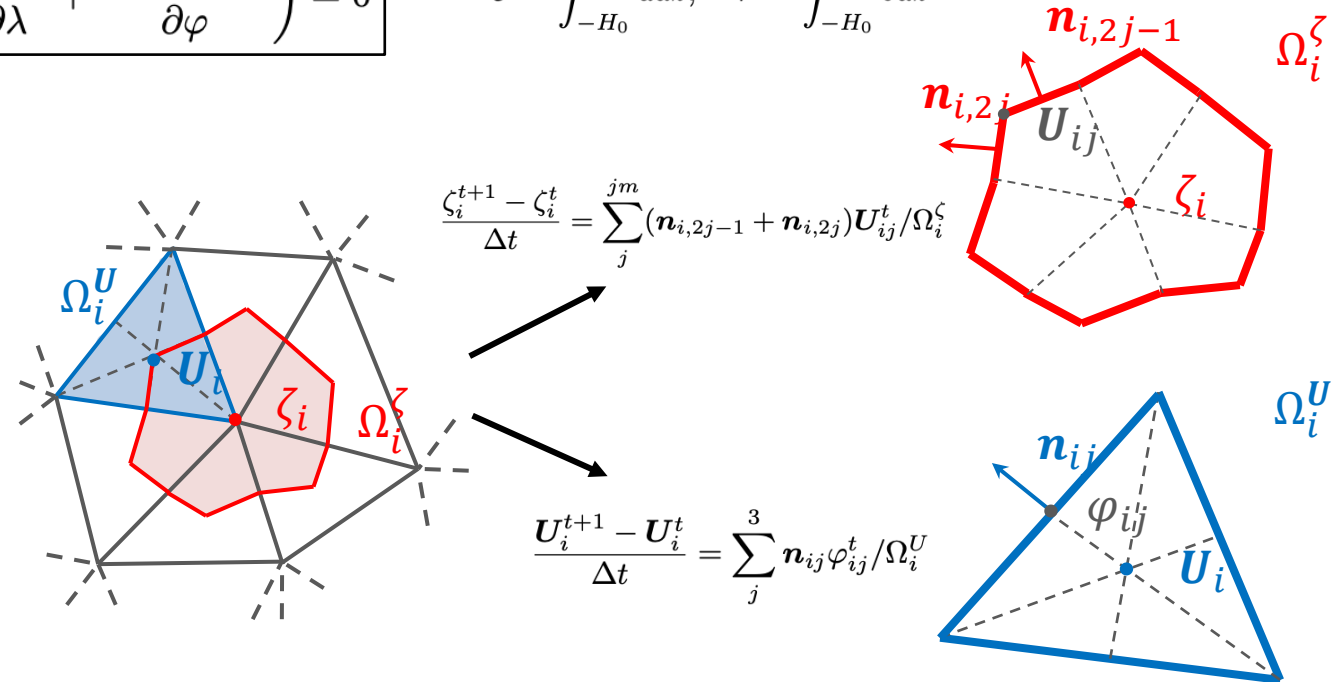
continuity equation

$$\frac{\partial \zeta}{\partial t} + \frac{1}{r \cos \varphi} \left( \frac{\partial uH}{\partial \lambda} + \frac{\partial vH \cos \varphi}{\partial \varphi} \right) = 0$$

$$U \equiv \int_{-H_0}^{\zeta} u dz, \quad V \equiv \int_{-H_0}^{\zeta} v dz$$

## ➤ Arakawa B-grid configuration

- current speed flux (vector)
  - placed at centroid of triangles
  - control volume is  $\Omega_i^U$
- sea-surface height (scalar)
  - placed at nodes
  - control volume is  $\Omega_i^\zeta$



# Numerical scheme

- The following scheme is adopted.
  - The Euler forward-backward scheme is adopted for a time integration.
  - The biharmonic filter is implemented to remove grid-scale noise.  
*cf. FESOM2 (Danilov et al, 2017)*
  - The implicit gravity-wave radiation condition is adopted for the open boundary condition.  
*cf. Chapman, 1985*
  - The first-order upwind differencing scheme is adopted for an advection term.
  - A parabolic drag coefficient is chosen.  
*cf. Peng & Li, 2015*

# Comparison of specification

	Previous model	New model
Model	two-dimensional linear model	two-dimensional nonlinear model
Grid	Lat-Lon, staggered grid (Arakawa C-Grid)	Unstructured (triangular) grid (Arakawa B-Grid)
Region	0 – 46°N, 95°E – 160°E	0 – 50°N, 95°E – 180°
Resolution	2' x 2' (~3.7 km)	1.5 km~50 km
Forecast period	72 hours	132 hours
Initial time	00, 06, 12, 18 UTC	00, 06, 12, 18 UTC
Ensemble member	no-typhoon case: 1 member (GSM) typhoon case: 6 members (GSM + 5 members chosen from GEPS)	no-typhoon case: 1 member (GSM) typhoon case: 52 members (GSM + GEPS 51 members)
Atmospheric forcing	GSM (~27km) GEPS (~55km)	GSM (~20 km) GEPS (~27 km)



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